

# Évènement de convection profonde de l'hiver 2011-2012 dans la mer d'Irminge: observation à l'échelle du bassin grâce aux données Argo

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# Summary

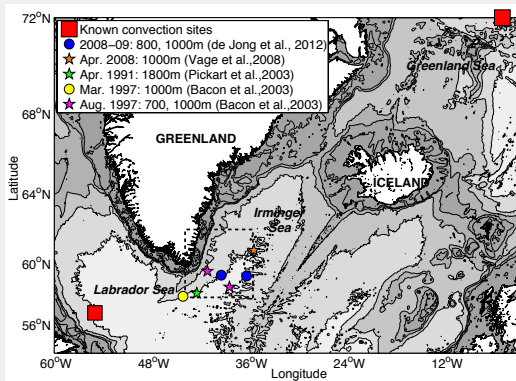
Context and Argo Data

The 2011-2012 winter deep convection event

Explanation by atmospheric forcings

Interannual variability

## Bibliographic context



*Convection sites in the North Atlantic Ocean and observations of deep mixed layers in the Irminger Sea*

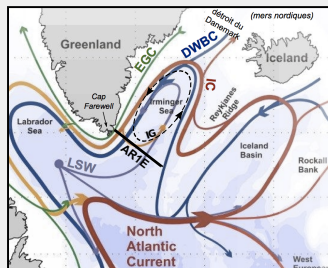
Past observations of convection in the Irminger Sea are **limited in space and time**

- ⌵ Lack of data, especially during wintertime
- ⌵ Focus on the Labrador site

# Mechanisms of convection in the Irminger Sea

## 1. Preconditioning

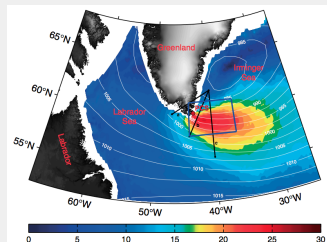
- Cyclonic circulation : Irminger Gyre [Lavender et al., 2000]
- Isopycnal doming [Pickart et al., 2003]
- LSW in the Irminger intermediate layers [Pickart et al., 2003]



North Atlantic circulation [P. Lherminier, LPO]

## 2. Atmospheric forcings : Greenland Tip Jets

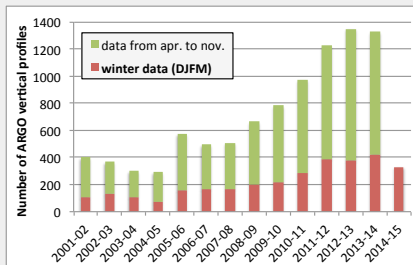
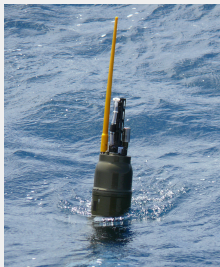
- regional-scale atmospheric events
- high wind speed (westerly)
- duration < 1 day
- induce heat loss ( $1000 \text{ W} \cdot \text{m}^{-2}$  locally)



Mean (1999-2002) Tip Jet QuikSCAT winds ( $\text{m} \cdot \text{s}^{-1}$ )

[Vage et al., 2009]; Blue box : 'TJ box'

## Argo profiles in the Irminger Sea



## Between Sept. 2011 and Sept. 2012 :

- 1209 vertical profiles from 253 different floats
- 4 floats have particularly been studied (4901163, 4901165, 4901166 and 5902298)
- 1 float with an  $O_2$  sensor during convection

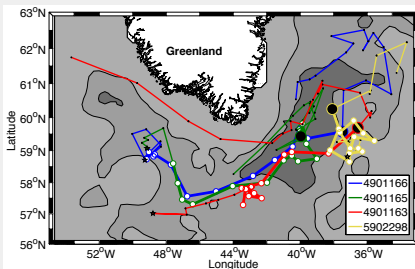
# Questions

The Irminger Sea ARGO sampling during the 2011-2012 winter, 3 to 4 times greater than for preceding winters, permit to identify a deep convection event. For this winter, thanks to the ARGO data, we now have informations about the ocean AT BASIN SCALE, DURING the convection event.



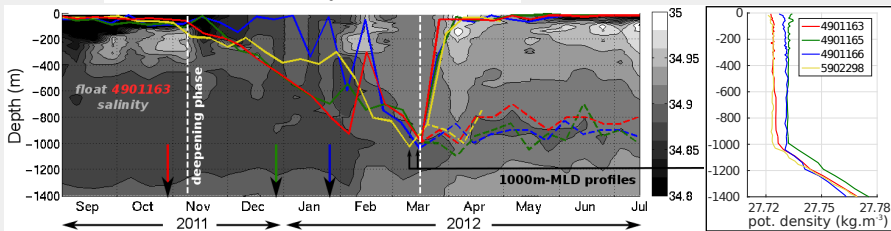
- ⌘ What is the spatial extent of deep convection in the Irminger Sea ?
- ⌘ Can we identify the sequence of atmospheric forcings responsible for convection ?
- ⌘ Can we use the 2011-2012 event to have a better understanding of the past events that could not be observed because of the lack of data ?

# Deepening of the mixed layers along floats trajectories

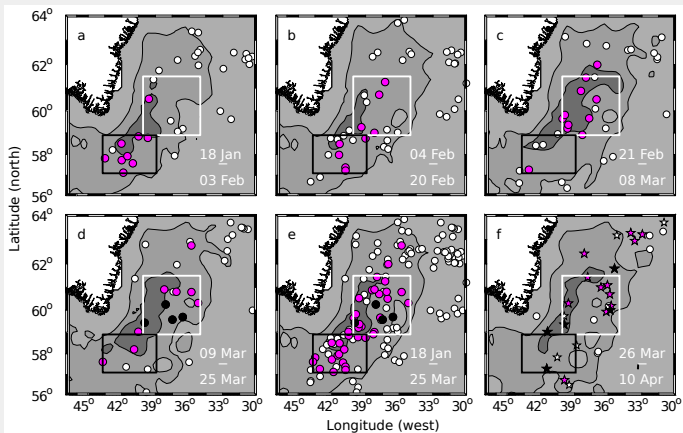


*WHITE dots :*  
floats positions during the  
deepening phase

*BLACK dots :*  
1000m-MLD positions



## Spatial extent of the convection



*Symbols : MLD (white : < 680m - magenta :  $\geq$  680m - black :  $\approx$  1000m)*

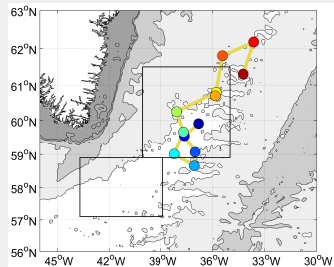
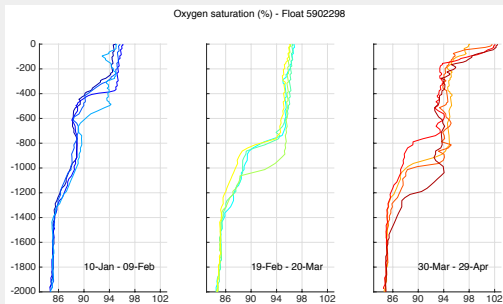
*Shade : dynamic topography (contours : -65, -55cm)*

*Boxes : convection areas (white : north box - black : south box)*





## Local formation of the convection

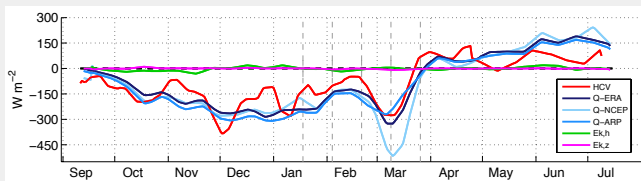


Oxygen data of 5902298 float show that convection occurs **locally** in the Irminger Sea

## Heat budget along floats trajectories

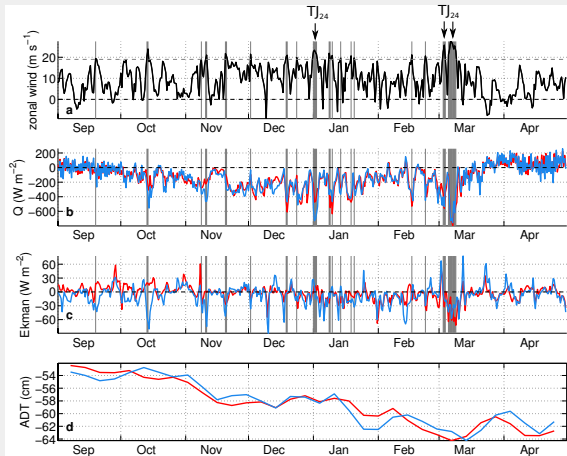
## Equation of heat budget [de Boissésion et al., 2010]

$$h\delta_t \langle T \rangle = \frac{F_{net}}{\rho_0 C_p} - U_{Ek} \delta_x T - V_{Ek} \delta_y T - [\langle T \rangle - T(-h)] w_{Ek}$$



- ⤵ Air-sea heat flux explain most of the mixed layer heat variation
- ⤵ Ekman terms are lower by one order of magnitude than air-sea heat flux
- ⤵ Heat losses are exceptionally strong early March

# Winter 2011-2012 Greenland Tip Jets



*Black curves : TJ box*

*Red curves : north box*

*Blue curves : south box*

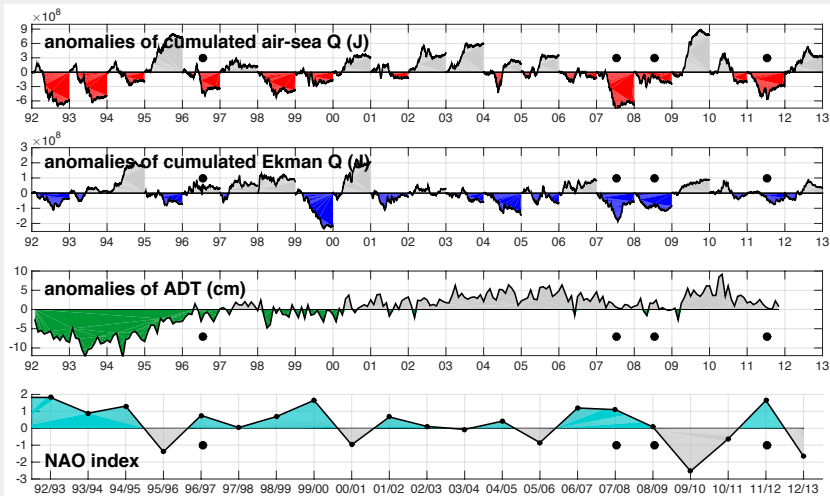
## Tip Jet criterion

zonal wind  $> 19 m.s^{-1}$

## Winter 2011-2012 inventory :

- Total : 18 TJ (Sep. - Mar.)
- 3 longer than a day
- 2 very exceptional early March

## Back to the early 1990s...



## Conclusions

- Thanks to ARGO data, it is the first observation of Irminger Sea deep convection, **at basin scale** and **during the whole period of convection**
- Convection zone located east and south of Cap Farewell
- Deepening (2 months) : firstly progressive, then interrupted (Feb.), and finally strongly restarted bringing ML at 1000m. Fast restratification
- Heat budget along floats trajectories : air-sea  $Q$  are mainly responsible for the ML heat content variation
- Late intense TJ early March deepened the MLs up to 1000m. Without these late TJ, MLs would probably not reach 1000m and convection would have stopped in February : so, finally a local small-scale atmospheric event has influence on a larger scale oceanic event
- The effect of Ekman advection is not negligible : its contribution can reach 10% of the total heat loss
- NAO, Ekman,  $Q$ , cyclonic circulation : indicators of deep convection for the Irminger Sea, for past years that have no observation

Reference : **Piron et al., 2015 (submitted in DSR1)**

THANK YOU !

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